ISOTOPE RATIOS IN THE MARTIAN UPPER **Atmosphere Measured by** MAVEN NGIMS

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NTRODUCTION

- The processes responsible for atmospheric escape to space enrich the atmosphere in the heavier isotopes of an element, because the lighter isotopes escape at a faster rate. This is true for two reasons:
 - Less energy is required for a lighter isotope to escape.
 - The upper atmosphere, from which escape occurs, is enriched in the lighter isotopes due to diffusive separation.

- The mass spectrometer on MAVEN, called NGIMS, measures the abundances of neutral and ionic species, and their isotopologues.
- We use the NGIMS data to investigate the upper atmospheric C and O isotope ratios in CO₂, including their vertical and horizontal variation. The variation of these and other isotope ratios is important for understanding the population of escaping atoms.



• LIGHT ISOTOPE • HEAVY ISOTOPE

NTRODUCTION

- To obtain an isotope ratio from the NGIMS data, we simply take the ratio of two *m/z* channels.
- We can then calculate a δ value relative to a standard, e.g. Vienna Pee Dee Belemnite or Standard Mean Ocean Water:

$$\delta^{13} \mathbf{C} = \begin{bmatrix} \left(\frac{{}^{13}\mathbf{C}}{{}^{12}\mathbf{C}}\right)_{\text{sample}} \\ \hline \left(\frac{{}^{13}\mathbf{C}}{{}^{12}\mathbf{C}}\right)_{\text{standard}} - 1 \end{bmatrix} \times 1000 \%$$

- Mean profiles are produced from bins of sequential orbits to give longitudinal averages.
- CO₂ density is used as the vertical coordinate because it is a convenient proxy for total atmospheric density or pressure that is directly measured by NGIMS.



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MEAN PROFILES

- Shown here are:
 - $\delta^{13}C_{VPDB}$ in CO₂ from m/z channels

45 (¹³C¹⁶O₂) and 44 (¹²C¹⁶O₂)

– δ¹⁸O_{vsmow} in CO₂ from m/z channels

46 (${}^{18}O^{12}C^{16}O$) and 44 (${}^{16}O^{12}C^{16}O$)

at L_s ~ 235° (MY 33), 3–5 p.m. local time.

• In blue is the slope and magnitude of the isotope ratio we would expect given the masses of the two species and the density of the lighter isotopologue.



- We also bin these isotope ratios on variables like Martian local time, latitude, and season, as we have for the neutral species (Stone et al. 2022, JGR: Planets).
- Here, data from the entire mission is binned on altitude and local time.
 - Because there is a significant temperature variation between the dayside and the nightside, the expected diurnal variation due to subsolar to antisolar transport is convolved with the contraction of the atmosphere on the nightside.



 Due to variation of the temperature, the mixing ratios of the two isotopologues change across a constant altitude surface simply because of the change in scale height of the atmosphere.

$$H = kT/mg$$



- When CO₂ density is used as the vertical coordinate, the effect of this contraction of the atmosphere on the nightside is not present and the diurnal variation due to transport is more clear.
- We find substantial variation of δ¹³C in CO₂ over a Martian day.
- This is the most significant trend observed in the variation of the isotope ratios. It is observed across all seasons and latitudes.



- A similar diurnal variation is observed for δ¹⁸O in CO₂.
- The diurnal variation of an isotope ratio in the upper atmosphere has important implications for the isotopic composition of the gas which can escape to space.

 In this case, the destruction of CO₂ leads to the formation of O₂⁺, then the escape of O via dissociative recombination.





- Vertical profiles of isotope ratios in CO₂ have been obtained from the NGIMS data.
- The isotope ratios observed in CO₂ vary substantially with local time.
- Measurement of these ratios their variation in the upper atmosphere is important to characterize the population of escaping atoms.
- These measurements are obtained in the NGIMS neutral mode, but we can also look at isotope ratios in the ionosphere.
 - For example, the isotope ratios in O_2^+ .
- We will explore the variation of the isotope ratios with respect to latitude and season and amid dust storms.
- Further analysis and detailed uncertainty calculations are required.